Palynology of Kamthi Formation in Godavari Graben

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Five palynological assemblages have been recognised in the Kamthi Formation of Godavari Graben. The assemblages are characterised by overall abundance of striate-disaccate pollen. The associated palynofossils, however, distinguish the palynoassemblages at various levels of the formation. The Lower Member containing coal seams shows unequivocal resemblance with the Raniganj palynoflora of Damodar Valley and Son-Mahanadi coalfields. Distinct change in lithology at the beginning of the Middle Member is marked by the appearance of *Parasaccites* rich assemblage simulating a cooling phase akin to that observed during the Talchir period. The younger sediments of the Middle Member contain *Corisaccites-Guttulapollenites* Assemblage and *Densipollenites* Assemblage in order of succession representing the uppermost Permian palynoflora in Godavari Graben. The youngest assemblage also indicates a close proximity to the Permian-Triassic transition, thus making the Kamthi Formation a time transgressive unit.

Key-words—Palynology, Kamthi Formation, Raniganj Formation, Permian/Triassic Boundary, Godavari Graben (India).

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साराँश

गोदावरी द्रोणिका में कामथी शैल-समृह का परागाणविक अध्ययन

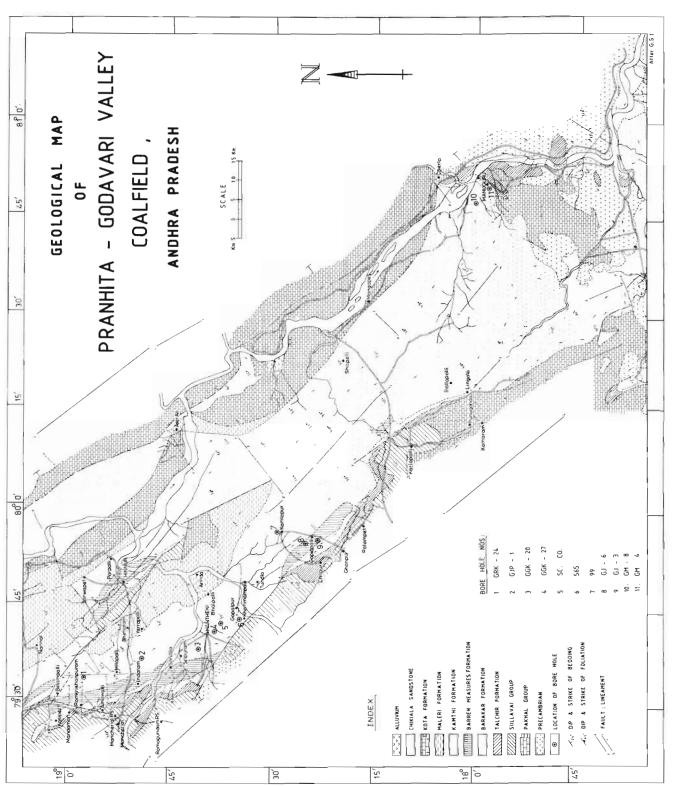
म्रेश चन्द्र श्रीवास्तव एवं नीरजा झा

गोदावरी द्रोणिका के कामथी शैल-समूह में पाँच परागाणिवक समुच्चय बनाये गये हैं। ये समुच्चय रेखीय-द्विकोष्ठीय परागकणों की बाहुत्यता से अभिलक्षणित हैं। सहयुक्त अश्मित परागकण शैल-समूह के विभिन्न स्तरों पर विद्यमान परागाणिवक समुच्चयों को विभाजित करते हैं। कोयला-सीमों से युक्त अधिर सदस्य सोन-महानदी कोयला-क्षेत्रों एवं दामोदर घाटी के रानीगंज परागाणुवनस्पतिजात से स्पष्ट समानता व्यक्त करता है। मध्य सदस्य के प्रारंभ में शैल-विन्यास में सुव्यक्त परिवर्तन परासेक्काइटिस से प्रभावी समुच्चय की उपिन्थित का द्योतक है जो कि तालिचर काल में प्रेक्षित शीतलन अवस्था से हर प्रकार से मिलता है। मध्य सदस्य के अल्पायु वाले अवसादों में कोरिसेक्काइटिस-प्रदुलापोलिनाइटिस समुच्चय तथा डेन्सीपोलिनाइटिस समुच्चय आरोही क्रम में विद्यमान हैं जिससे गोदावरी द्रोणिका में उपिरतम् परमी युगीन वनस्पतिजात की उपिन्थित इंगित होती है। अल्पतम् आयु वाली समुच्चय भी परमी-त्रिसंघी परिवर्तन से घनिष्ठता इंगित करती है जिससे कि कामथी शैल-समूह का स्वरूप एक उत्क्रमी इकाई के रूप में सामने उभर कर आता है।

SEDIMENTS overlying the Barakar Formation and underlying the Maleri Formation in the Godavari Graben were earlier included within the Kamthi ormation (King, 1881). After the identification of Barren-Measures Formation in Bheemaram area (Sengupta, 1970) and Ramagundam area (Ramanamurty, 1979), the strata occurring between

the Barren-Measures and Maleri Formation are referred to the Kamthi Formation (sensu Ramanamurty, 1985).

KamthiFormation is spread over a vast tract of Godavari Graben (Map 1) and exhibits significant lateral and vertical variation in lithofacies. The formation attains maximum thickness around



Map 1-Geological map of Pranhita-Godavari Graben, Andhra Pradesh.

Ramagundam and Mantheni areas. It has been divided into: (i) the Lower Member, consisting of greyish white, medium-grained, calcareous sandstones and a few coal seams; (ii) the Middle Member, marked by alternating sequence of medium-grained grey white sandstones, shales and variegated clays; and (iii) the Upper Member, comprising coarse-grained arenaceous facies incorporating coarse-grained ferruginous sandstones and brick-red siltstones. This member forms important topographic features in the Godavari Graben. The contact between the underlying Barren Measures and Lower Member of KamthiFormation is gradational and so also is the case between Lower and Middle members. The Middle Member is totally devoid of coal seams and the sandstone and shales bear a greenish tint and underlies the Upper Member with an unconformity.

The samples investigated have been obtained from following areas (Map 1).

1.	Ramkrishnapuram	—В. Н. no. GRK-24
2.	Jaipuram	—В. Н. no. GJP-1
3.	Mantheni	—S. C. Co. bore hole
4.	Khammampalli	—В. Н. no. 565
5.	Bhopalpalli	—В. Н. no. GJ-6
6.	Kamalpur	—В. Н. no. 99
7.	Manuguru	—B. H. nos. GM-4 and
		CM-8

In addition, data from following three boreholes have been incorporated (Bharadwaj *et al.*, 1987).

- B. H. no. GGK-20 Ramagundam-Mantheni area
 B. H. no. GGK-27
- 3. B. H. no. GJ-3 Chelpur area

PALYNOLOGICAL ASSEMBLAGES

The palynological investigations have led to recognition of five assemblages based on their morphographic characters and also quantitative representation. The distribution of various palynotaxa within the Kamthi Formation has been shown in Text-figure 1.

Assemblage 1

Faunipollenites-Striatopodocarpites Assemblage—This assemblage is characterised by dominance of striate-disaccate pollen chiefly represented by Faunipollenites and Striatopodocarpites. Scheuringipollenites, a nonstriate-disaccate, is a subdominant element. The presence of some other striate-disaccate genera is significant, viz., Verticipollenites, Labirites, Hindipollenites, Crescentipollenites, Distriatites, etc. Taeniate pollen Lunatisporites, Hamiapollenites, Corisaccites,

Lueckisporites occur in rare amounts and their behaviour is inconsistent. Nonstriate-disaccate pollen Alisporites, Falcisporites, Vitreisporites and Chordasporites also behave like taeniate-disaccate pollen grains. The percentage of trilete spores is comparatively low but the occurrence is fairly consistent. They are represented by Horriditriletes, Brevitriletes, Verrucosisporites, Gondisporites, etc.

The Faunipollenites-Striatopodocarpites Assemblage is lithologically associated with the Lower Member and has been recorded in bore-holes GRK-24, GGK-20, GGK-27, GJ-3, GJ-6, GM-4 and GM-8 (Text-fig. 2).

Assemblage 2

Faunipollenites-Striasulcites Assemblage—The dominance of striate-disaccate pollen remains similar as in the preceding assemblage. However, Scheuringipollenites gives way to the genus Striasulcites which is restricted to the present assemblage only. Further, Densipollenites also appears simultaneously but remains low in percentage. Falcisporites, Vitreisporites, Vesicaspora, etc. remain impersistent. Though the percentage of trilete spores declines, yet the occurrence of Gondisporites, Polypodiidites, Osmundacidites, etc. is significant.

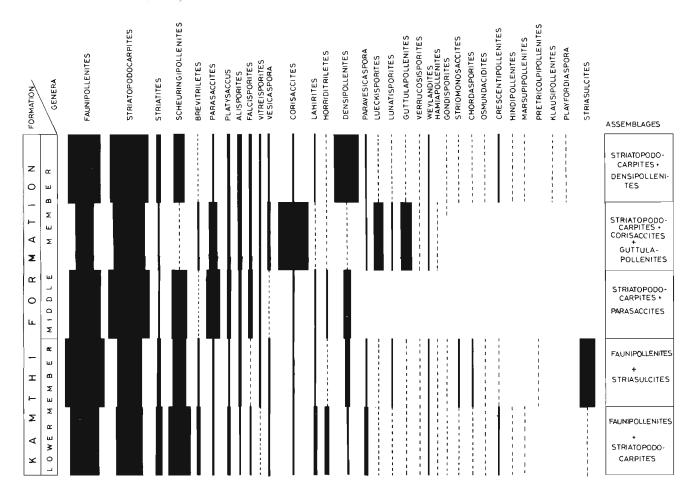
The Faunipollenites-Striasulcites Assemblage is well-developed in bore-hole GGK-27 between 580-385 m which marks the lower part of the Middle Member (Text-fig. 2), in bore-hole 565 between 202.25-203.75 m in carbonaceous clayey shale, in bore-hole GM-4 at 170 m in a grey clay overlying a coal band and in bore-hole GM-8 at 95-84 m.

Assemblage 3

Striatopodocarpites-Parasaccites Assemblage— In bore-hole GJ-6 (210-20 m) Striatopodocarpites maintains overall dominance and is followed by Faunipollenites (Text-fig. 2). Parasaccites assumes subdominance, next to the striate-disaccate pollen grains. Densipollenites, Corisaccites and Guttulapollenites occur in low percentages. The other taeniate disaccate genera are not as well represented. The trilete spores are rare. Falcisporites, Vitreisporites, etc. occur only sporadically. This assemblage is associated with sandy black shale, grey black shale and green sandstones and thus, occurs at the end of the coal-bearing sequence.

Assemblage 4

Striatopodocarpites - Corisaccites - Guttulapollenites Assemblage—The dominance of striate-disaccate pollen continues unabatedly in younger sediments (Middle Member). Densipollenites tends to increase comparatively but remains low.



Text-figure 1—Succession of various palynological assemblages in Godavari Grabven.

Corisaccites and Guttulapollenites together increase to subdominance, the former being more in number. Lueckisporites is also represented in significant percentages. Lunatisporites, Falcisporites, Vitreisporites appear regularly but in low percentages. In bore-hole GJP-1 (Text-fig. 2) Corisaccites reaches its maximum at 276 m and is associated with Guttulapollenites and Parasaccites. Towards north, in bore-hole GRK-24, this assemblage continues between 121-4-127.00 m. In bore-hole GGK-27, Corisaccites occurs between 334-241 m but is in low amounts while Guttulapollenites is almost absent. In other areas this assemblage is very poorly represented.

Assemblage 5

Striatopodocarpites Densipollenites Assemblage—This assemblage is fairly well-represented in almost all the areas investigated. The sediments overlying the Striatopodocarpites Corisaccites Guttulapollenites Assemblage are characterised by a significant rise in the representation of the genus Densipollenites.

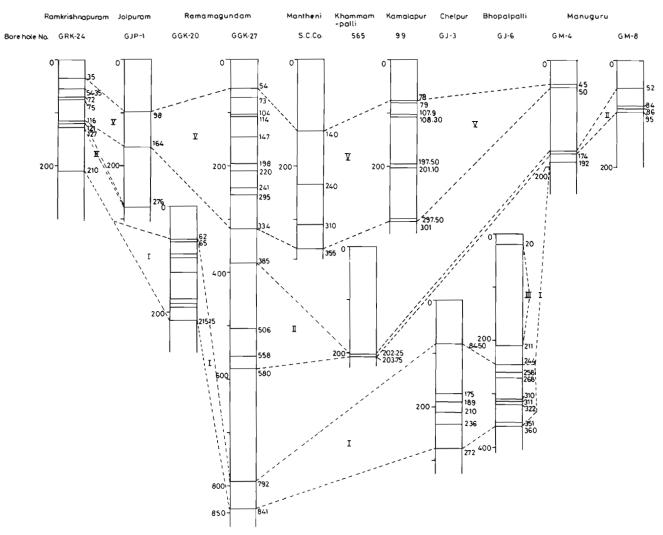
This assemblage succeeds *Striatopodocarpites-Corisaccites* Assemblage in bore-hole GRK-24 (116-35.95 m) and bore-hole GJP-1 (164.98 m). *Weylandites* shows a rising trend in bore-hole GRK-24. In bore-hole GGK-27 (222-54 m) *Densipollenites* records a perfect epibole and reaches overall dominance between 198-147 m. In other areas also *Densipollenites* has been observed to be closely associated with striate-disaccate pollen.

COMPARISON

A number of palynofloras described from other Lower Gondwana coalfields of India resemble one level or the other in the Kamthi Formation of Godavari Graben and their related stratigraphic comparability has also been summarised in Table 1.

Assemblage 1 closely compares with the palynological succession in Raniganj Formation of Damodar Valley (Bharadwaj & Tiwari, 1977; Bharadwaj, Tiwari & Anand-Prakash, 1979). The overall dominance of striate-disaccate genera continues throughout the entire succession.





Text-figure 2—Palynological correlation of assemblages in different bore-holes of Godavari Graben. V *Striatopodocarpites + Densipollenites*; IV *Striatopodocarpites + Corisaccites + Guttulapollenites*; III *Striatopodocarpites + Parasaccites*; II *Faunipollenites + Striatopodocarpites + Striatopodocarpites*.

However, *Indospora* and *Spinosporites* present, though in rare amounts, in Raniganj sediments of Damodar Valley have not been observed so far in the Kamthi sediments.

The palynoflora of the Jhingurdah Seam in Singrauli Coalfield (Tiwari & Srivastava, 1984) also compares closely in view of preponderance of striate-disaccate pollen grains. The trilete spores are, however, more frequent in Jhingurdah Seam which also has *Indospora* and *Kendosporites*.

The palynoflora of Raniganj Formation in Auranga Coalfield (Lele & Srivastava, 1979) has a preponderance of striate-disaccate pollen grains, but the percentage of trilete spores is relatively more. urther, *Mahudapollenites* and *Mammialetes* are absent in Assemblage 1.

An assemblage comparable to Faunipollenites-Striasulcites Assemblage is not known from any other basin except for the Striasulcites rich assemblage from Koel River Section in Hutar Coalfield (Shukla, 1983) which also has Potonieisporites, Scheuringipollenites and Faunipollenites. Potonieisporites is not present in this palynoassemblage of Godavari Graben. In addition, the genera Gondisporites, Polypodiidites, Osmundacidites, Corisaccites, Guttulapollenites, Alisporites, Vitreisporites, Falcisporites, Weylandites that occur regularly though in low percentages are absent in Hutar palynoassemblage.

Striatopodocarpites-Parasaccites Assemblage is also being first recorded from the Godavari Graben. In the Raniganj Formation of Damodar Valley, Parasaccites is present in Striatopodocarpites-Densipollenites Assemblage (Bore-hole RAD-5; Tiwari & Singh, 1983) but the percentage of this genus remains low.

Striatopodocarpites-Corisaccites Assemblage is comparable with a similar assemblage reported from

Table 1-Stratigraphic relationship of different palynoassemblages in Permian basins of India

Auranga	Coalfield						Lele & Srivastava (1979) Sukri River- Carbonaceous shale- Faunipollenites + Brevitriletes
ALLEY	Gopad River		Maheshwari (1977) Carbonaceous Shale-Striato- podocarpites + Densipollenites				
SON VALLEY	Singrauli Coalfield						Tiwari & Srivastava (1984) Jhingurdah Seam Faunipollenites Striatopodocar- pites
Raimahal	Basin		Tiwari & Tripathi (1984) Lower part of Dubrajpur Formation				T IL 8
, tron	Karanpura Coalfield						Bharadwaj et al. (1979) Horam Sibadih Section- Carbonaceous Shale Striato- podocarpites + Faunipollenites
	Ondal Area						et al. Kar (1970) nnia NalaB.H. RE9.Carbona- bona- ceous Shale, carpites B.H.RN9.Carbona- tites ceous Shale Faunipollenites Striatopodocar- pites
COALFIELD	North of Damodar River		Bharadwaj et al. Bharadwaj et al. (1979) Machkanda (1979) Nonia Nala Jhor Section Section—Grey Shale, Section A—Grey Laminated Shale, Shale Striatopo Carbonaceous Shale docarpites Densi: Nonia Khal Section—Garbonaceous Shale Banerij & Maheshwari (1975) Nonia Nala Section—Grey Shale Striatopodocarpites + Densi: pites + Densi:				Bharadwaj et al. a(1979) Nonia Nalz Section-Carbona- ceous Shale, Coal Striatopodocarpites Faunipollenites
RANIGANJ	South of Damodar River		Bharadwaj et al. (1979) Machkano Jhor Section Section Section Shale Striatopo-docarpites Densipollenites				Bharadwaj et al. (1979) Machhkand Jhor Section A—Carbonaceous Shale, Coal Section B—Carbonaceous Shale, Khaki Shale Striatopodocarpites* Faunipollenites
20026	7,643	sa	Densipollenites	Corisaccites Guttulapollenites	Parasaccites	Striasulcites	Striatopodocar. pites
		Assemblages	. BIVIOBODOCVBBILES	LS	E 2	ΤI	EVONIBOLLEN

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Table 1 (Contd.)

Lesser	Tethvan	Satoura			GODAVAR	GODAVARI GRABEN (Present investigation)	investigation)		
Himalaya	Himalaya	Basin	Ramkrishanpuram	Ramagundam	Chelpur	Bhopalpalli	Kamalapur	Khammampalli	Manuguru
Tiwari & Kumar (1986) Amkhal For- mation Striatopo- docarpites + Densipollenites	Tiwari & Tiwari et al. Kumar (1986) (1984) Malla Amkhal For- Johar, Kuling mation Shale Striatopo- Striatopo- docarpites + docarpites + Densipollenties	Bharadwaj et al. (1978) Bijori Formation Striatopodo. carpites + Densipollenites	Striatopodo. carpites + Densipolle. nites	Striatopodo. carpites + Densipolle. nites		Striatopodo. carpites + Densipolle. nites	Striatopodo- carpites + Densipolle- nites		Striatopodo. carpites + Densipolle. nites
		Striatopodo. Striatopodo- carpites + carpites + Guttulapollenites + Corisaccites + Corisaccites Guttulapoll:	Striatopodo. carpites + Corisaccites + Guttulapoll. enites	Striatopodo- carpites + Corisaccites + Guttulapoll- enties					
						Striatopodo. carpites + Parasaccites			
				Faunipollenites + Striasulcites				Faunipollenites + Striatopodocar- pites + Striasulcites	Faunipollenites + Striasulcites
			Faunipollenites Striatopodo carpites	Faunipollenites Striatopodo. carpites	Faunipollenites Striatopodo- carpites	Faunipollenites Faunipollenites + Striatopodo: Striatopodo: carpites carpites			Faunipollenites + Striatopodo curpites

1.55	DAMODAR VALLEY	G 0	D A V		٧	Α	RI		G	R	Α	В		Ε 1	N
AGE	G.S.I. (1977)	King (1881)	Ser	ngupta (1970)		Kutty	et. al. (1987)	G.	S.I. (1882)		Raiv	verman	et a	l. (198	1985)
TRIASSIC	PANCHET	U	I P I	P E	1	R	៤	0 I	N	D	W	Α	N	А	
TRI		К.	K A M	UPPER MEMBER MIDDLE		K A M	UPPER MEMBER MIDDLE	K	UPPER MEMBER		KUDUREPALLI / (HINTALPUDI MANER				
		Α	T H I	MEMBER LOWER MEMBER	-	T H	MEMBER LOWER MEMBER	M							
z		M T		ON STONE SHALES		I N F	LITHOZONE - 4	Т	MIDDLE		KHANAPUR				
P E R M I A	RANIGANJ					RA	LITHOZONE - 3	н	MEMBER		JAIPURAM				
		Н	DADAKAD			KAMTHI	LITHOZONE - 2	ı	LOWER MEMBER REN MEASURES		POTAMADUGU / BALHARSHAH				
	BARREN MEASURES			BARAKAF			LITHOZONE - 1	BARI			BELLAMPALLI				
	BARAKAR	BARAKAR	KAR			ВА	BARAKAR		BARAKAR		BARAKAR				
	TALCHIR	TALCHIR	TALCHIR			TALCHIR		TALCHIR			TALCHIR				

Table 2-Stratigraphic status of Kamthi Formation in Godavari Graben

the Bijori (Sukhtawa) Formation in Satpura Basin (Bharadwaj et al., 1978). Corisaccites-Guttulapollenites occupy a third place in the order of dominance in Satpura Basin but in Godavari Graben this pollen-complex assumes overall dominance.

Striatopodocarpites, Densipollenites Assemblage is fairly well distributed in Damodar Valley, Singrauli and Auranga coalfields and Satpura Basin. In Damodar Valley this assemblage is present in the Late Permian (Tiwari & Singh, 1983) and continues up to the base of the Lower Triassic. In Satpura Basin, Densipollenites rises to dominance near the top of the Bijori Formation succeeding the Corisaccites rich assemblage (Bharadwaj et al., 1978).

The striate-disaccate rich assemblage associated with *Densipollenites* from Gopad River Section (Maheshwari, 1967) resembles Assemblage 5. The palynoassemblages described from the lower part of the DubrajpurFormation in Rajmahal Basin (Tiwari & Tripathi, 1984), the Kuling Shale in the Malla Johar area, Kumaon Himalaya (Tiwari *et al.*, 1984), and the Amkhal Formation in the Lesser Himalaya of Tehri Garhwal District (Tiwari & Kumar, 1986) contain a large number of taxa found in *Striatopodocarpites-Densipollenites* Assemblage and thus bear a qualitative resemblance.

DISCUSSION

Raiverman, Rao and Pal (1985) have reviewed the entire geological succession in PranhitaGodavari Graben and have provided a new classification on the basis of photogeological techniques and detailed geological field mapping. They have raised the status of Kamthi to a group and divided it into five formations. Their correlation with earlier classification schemes is shown in Table 2. Kutty et al. (1988) have restricted the limits of the KamthiFormation as equivalent to the Upper Member of the Kamthi Formation only. However, in the present investigation the classification in vogue by the Geological Survey of India (Raja Rao, 1982) has been followed. The non-coal-bearing predominantly arenaceous sequence of Barren Measures conformably grades into the coal-bearing Lower Kamthi Member. However, the palynological assemblage changes from the Densipollenites rich assemblage of Barren Measures to Faunipollenites-Striatopodocarpites rich assemblage with near absence of Densipollenites of Lower Kamthi Member. The appearance of Gondisporites, Falcisporites, Vitreisporites, Corisaccites, Lunatisporites, Weylandites, etc. though in rare amounts, nevertheless indicates a younger affinity being characteristic of the Raniganj and Panchet assemblages in Damodar Valley. Thus, the palynoflora of the lower part of the Raniganj formation appears almost coeval with that of the Lower Member of Kamthi Formation. The Faunipollenites-Striasulcites Assemblage is associated with coal seams in Manuguru area, but in Ramagundam area it is recorded above the coalbearing horizon in the beginning of Middle Member.

The sediments between 792-580 m in bore-hole GRK-27, which include the transition from Lower to Middle Member, have not yielded pollen and spores hence it is difficult to comment on the exact stratigraphic status of Striasulcites but from the present evidences it appears that the Striasulcites assemblage reached its peak at the end of coalforming phase. The sedimentation of the Lower Member appears to have commenced in a prevailing fluvial environment in relatively deep and narrow to straight sinuous channels. A warm and humid climate seems to have provided a luxurient growth of gymnosperms and pteridophyte which have contributed largely to the coal formation.

The incoming of *Parasaccites* at the base of the Middle Member coupled with the greenish sandstone and shales is significant as it marks the culmination of coal-forming phase. Parasaccites is known to have been associated with glacial and cold climate in the Lower Gondwana Sequence and hence a similar cooling though short-lived is envisaged to have been the cause of culmination of coal forming process in the Godavari Graben. Densipollenites commencing after the Parasaccites phase, continued up to the middle part of the Middle Member which is comparable with the coal-less Barren Measures, denoting a comparative aridity in the climate.

Striatopodocarpites-Densipollenites Assemblage continues up to the top of the Ranigani Formation (Upper Permian) and is succeeded by the cavate-cingulate-taeniate rich assemblage of the Panchet Formation (Lower Triassic, Tiwari & Singh, 1983). In the Godavari Graben epibolic development of *Striatopodocarpites-Densipollenites* Assemblage is noticed within the lower 334 m of the Middle Member in Ramagundam area. The decline of Bharadwaj, D. C., Srivastava, S. C., Ramanamurty, B. V. & Jha, N. Densipollenites is recorded at the top of the sequence indicating a close proximity towards the Lower Triassic boundary. There is every likelihood that the Permian-Triassic boundary lies within the upper part of the Middle Member. Lithologically also the sandstones and shales of this Member are distinct from the underlying Lower Member being greenish in colour and devoid of coal. Some palegreen clays, apparently looking like Talchir needle shales (King, 1881) are exposed around Sitampet. Similar shales are also exposed near Rangaipalli and Jillapalli villages in shallow wells. These khaki green shales which are apparently younger than the Middle Member of Ramagundam-Mantheni area may infact prove to be equivalent to the Panchet Formation.

Evidences indicate that the Lower Member containing coal seams is equivalent to the lower part of the Ranigani Formation of Damodar Valley while the lower part of the Middle Member corresponds to the upper part of the Raniganj Formation and the Bijori Formation in Satpura Basin. The Lower Triassic equivalent palynoflora is expected within the upper part of the unexplored Middle Member. In other words, the Kamthi Formation is time-transgressive palynologically, ranging in age from Upper Permian to ?Triassic.

CONCLUSION

Five palynological assemblages can be recognised in the Kamthi Formation of the Godavari Graben. Assemblage 1 from the Lower Member shows a uniform similarity with Ranigani palynoassemblage of Damodar and Son-Mahanadi grabens. This assemblage is associated with the coalbearing sediments of the Ranigani Formation. The lithological changes at the base of Middle Member are associated with the appearance of Striatopodocarpites-Parasaccites rich assemblage (Assemblage 3). These changes together document an intensive cooling phase, though restricted, simulating the one observed during Talchir Formation. The younger palynological assemblages (Assemblages 4 and 5) are more closer to the palynological succession observed in Satpura Basin and Salt Range, Pakistan. Densipollenites rich assemblage in Godavari Graben declines within the Middle Kamthi Member and a larger part of the sequence yet remains to be explored palynologically. The Permian-Triassic boundary is expected within the Middle Kamthi Member.

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