

# Palynology and palaeoenvironment of Late Permian Sawang OCM, East Bokaro Coalfield, Damodar Basin, India

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## ABSTRACT

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Palynodating of Sawang Open Cast Mine (OCM) section from East Bokaro Coalfield, Damodar Basin has been done. Recovered palynofossils are characterized by the dominance of *Striatopodocarpites*–*Faunipollenites* complex. The other stratigraphically significant taxa recorded from this section are *Gutulapollenites hannonicus*, *Crescentipollenites fuscus*, *Rhizomaspora indica*, *R. triassica*, *Distriatites* sp., *Weylandites lucifer*, *Microfoveolatispora gondwanensis*, *Dictyotriletes invisus*, *Indotriradites* spp., *Arcuatipollenites pellucidus*, *Alisporites indicus* and *Klausipollenites schaubergeri*. Other rare palynotaxa include *Parasaccites*, *Corrisaccites*, *Dicapipollenites*, *Striomonosaccites*, *Barakarites*, *Plicatipollenites*, *Scheuringipollenites*, *Densipollenites*, *Callumispora*, *Tiwariasporis*, *Praecolpatites* and *Distriatites*. On the basis of the total palynocomposition, the studied section has been dated as Late Permian age. This age correlation also gets support from comparative studies with similar palynoassemblages known from other coalfields of Indian Gondwana basins such as Damodar, Son–Mahanadi, Rajmahal, Wardha–Godavari and Satpura basins of India and from known Gondwanan continents. Palynofossil evidences indicate prevalence of warm and humid conditions. The dominance of conifers and subdominance of Glossopterids, cordaites and low percentage of triletes (filicopsids and lycopsids) suggests that the Inland Sawang OCM was deposited under freshwater environment.

**Key-words**—Palynofossils, Late Permian, Sawang Open Cast Mine, Palaeoenvironment, Bokaro Coalfield, Damodar Basin, Jharkhand.

## भारत की दामोदर द्रोणी के पूर्वी बोकारो कोयला क्षेत्र की अंतिम पर्मियन सवंग ओ सी एम का पुरागाणुविज्ञान एवं पुरापर्यावरण

श्रीकांत मूर्ति एवं ए. रजनीकांत

### सारांश

दामोदर द्रोणी के पूर्वी बोकारो कोयला क्षेत्र से प्राप्त सवंग विवृत खान (ओ सी एम) का पुरागाणु कालनिर्धारण कर दिया गया है। प्राप्त पुरागाणुविज्ञान जीवाश्म *स्ट्रिएटोपोडोकार्पाइटिस* – *फॉनीपोल्लेनाइटिस* सम्मिश्रण की प्रभुत्वता से अभिलक्षित है। इस खंड से अभिलिखित अन्य स्तरिक रूप से सार्थक टैक्सा *गुटुलापोल्लेनाइटिस हन्नोनिकस*, *क्रेसेंटिपोल्लेनाइटिस फस्कस*, *राइजोमास्पूरा इंडिका*, *आर. ट्रिएसिका*, *डिस्ट्रीआटाइटिस* जाति, *वीलंडीटाइटिस लुसीफर*, *माइक्रोफेवोलैटिसपोरा गोंडवानेन्सिस*, *डिक्टीओट्रिलेटीज इन्वीसस*, *इंडोट्रिरैडाइटिस* जातियां, *अर्कुएटिपोल्लेनाइटिस पेल्लुसाइडस*, *अलीस्पूराइटिस इंडिकस* और *क्लोसीपोल्लेनाइटिस स्कोबर्गेरी* हैं। अन्य दुर्लभ पुरागाणु टैक्सा में *पैरासैक्काइटिस*, *कोरिसैक्काइटिस*, *डिकपीपोल्लेनाइटिस*, *स्ट्रीमोनोसैक्काइटिस*, *बराकराइटिस*, *प्लिकेटिपोल्लेनाइटिस*, *स्युरिंगीपोल्लेनाइटिस*, *डेन्सिपोल्लेनाइटिस*, *केलुमिस्पूरा*, *तिवारीआस्पूरिस*, *प्रीकॉल्ल्याटाइटिस* एवं *डिस्ट्रीएटाइटिस* सन्निहित हैं। समग्र पुरागाणु संघटन के आधार पर, अध्ययन किया गया खंड अंतिम पर्मियन काल के रूप में कालनिर्धारण किया गया है। यह आयु सहसंबंध भारतीय गोंडवाना द्रोणियों जैसे कि दामोदर, सोन–महानदी, राजमहल, वर्धा–गोदावरी व भारत की सतपुड़ा द्रोणियों और ज्ञात गोंडवाना की अंतर्वस्तुओं के अन्य कोयला क्षेत्रों से ज्ञात सदृश पुरागाणुसमुच्चयों के साथ तुलनात्मक अध्ययनों से भी समर्थन प्राप्त करते हैं। पुरागाणुजीवाश्म प्रमाण कोष्ण एवं आर्द्र स्थितियों की व्यापकता इंगित करते हैं। शंकुवृक्षों की प्रभुत्वता और ग्लोसोप्टेरिड, कॉर्डैटीज की उपप्रभुत्वता तथा त्रिअरीय की अल्प प्रतिशतता (फिलिकॉप्सिड व लायकॉप्सिड) सुझाते हैं कि अंतर्देशीय सवंग ओ सी एम अलवणजल पर्यावरण के अधीन निक्षेपित हो गया था।

**सूचक शब्द**—पुरागाणुजीवाश्म, अंतिम पर्मियन, सवंग विवृत खान, पुरापर्यावरण, बोकारो कोयला क्षेत्र, दामोदर द्रोणी, झारखंड।

## INTRODUCTION

THE Damodar Basin is an important Indian coal basin and spreads in the Indian states of West Bengal and Jharkhand. Important coalfields in this basin are Raniganj, Jharia, East Bokaro, West Bokaro, Ramgarh, South and North Karanpura. The Bokaro Coalfield is situated in Hazaribagh and Giridih districts of Jharkhand State. The basin is an elongated strip of Gondwana sediments stretching over 64 km from east to west and 12 km in width. The name Bokaro was given by Williams (1846–47) after the Bokaro River which flows for a distance of about 40 km in this region. This coalfield has been divided into two distinct zones, East Bokaro Coalfield (EBC) and West Bokaro Coalfield (WBC) in the Lugu Hill massif (Dutt, 1944–51; Kumar & Sahay, 2001). Further, the part of the coalfield, east of longitudes  $85^{\circ}42'$  is commonly known as East Bokaro Coalfield (EBC) which is an area of about 237 km<sup>2</sup> and located between latitudes  $23^{\circ}44'$  &  $23^{\circ}49'$  N and longitudes  $85^{\circ}42'$  &  $86^{\circ}4'$  E in Damodar Basin (Raja Rao, 1987). The Sawang OCM situated in the north–western part of the EBC and its geographical location is marked by latitudes  $23^{\circ}47'40''$  to  $23^{\circ}48'28''$  N and longitudes  $85^{\circ}50'37''$  to  $85^{\circ}51'50''$  E and included in Survey of India Toposheet No.73E/13 (Pophare & Varade, 2004a, b). The coal–bearing sediments of Sawang area have unconformable contact with the underlying metamorphic rocks (Pophare *et al.*, 2008). The generalized stratigraphic succession of the East Bokaro Coalfield after Raja Rao (1987) is shown in Table 1.

The Indian Gondwana basins are rich in palynofossils (Tiwari, 1999). Various formations of these basins have yielded rich palynofloral assemblages representing lower, middle and upper Permian ages. The palynology of Damodar Basin is well established—(Bharadwaj, 1962; Bharadwaj &

Salujha, 1964; Bharadwaj & Srivastava, 1969; Kar, 1968; Tiwari & Singh, 1981; Tiwari & Tripathi, 1992; Srikanta Murthy *et al.*, 2010; Srikanta Murthy *et al.*, 2014).

Negligible palynological work has been done till now in the Bokaro Coalfield. Lele (1973) recovered two micro–floral assemblages belonging to the Talchir Formation from the Dudhi River section, West Bokaro Coalfield. An Olenekian palynoflora was recovered by Pal and Ghosh (1994) from the Panchet Formation exposed at the base of Lugu Hill from Bokaro Coalfield. Vijaya *et al.*, (2012) carried out a detailed palynological study from Borehole EBM–1 (East Bokaro Coalfield, Muditoli Block,) and suggested a Permian age for this borecore sequence. Recently, Srikanta Murthy *et al.*, (2016) carried out palynological and petrographical studies from Borehole EMB–2 (western part of East Bokaro Coalfield) and suggested Permian age.

The present investigation is focused on the palynology of the sediments located above the coal seam in Sawang OCM Section (Fig. 2A). Palynodating and correlation have been attempted, and palaeoenvironment has also been inferred. Comparisons with other Late Permian palynofossils of Gondwanan basins in Peninsular India, Australia, Antarctica, Africa and South America have also been done.

## GENERAL GEOLOGY OF THE COALFIELD

The Gondwana sediments in the EBC are represented by Talchir, Barakar (Early Permian), Barren Measures Raniganj (late Permian), Panchet (early Triassic) and Supra Panchet (middle to late Triassic) formations (Table 1, after Raja Rao, 1987). The scattered rocks of Talchir Formation unconformably overlie the basement rocks in the north–eastern periphery of the coalfield, and includes tillite, greenish

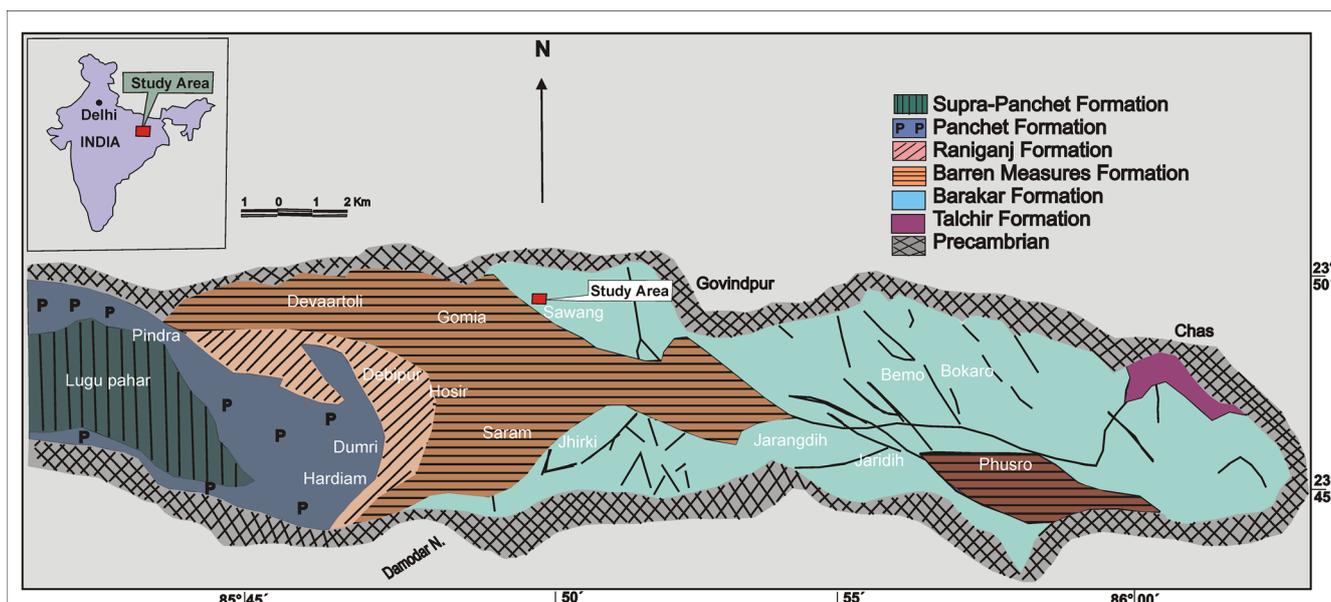


Fig. 1—Geological map of East Bokaro Coalfield showing the location of the study area (after Raja Rao, 1987).

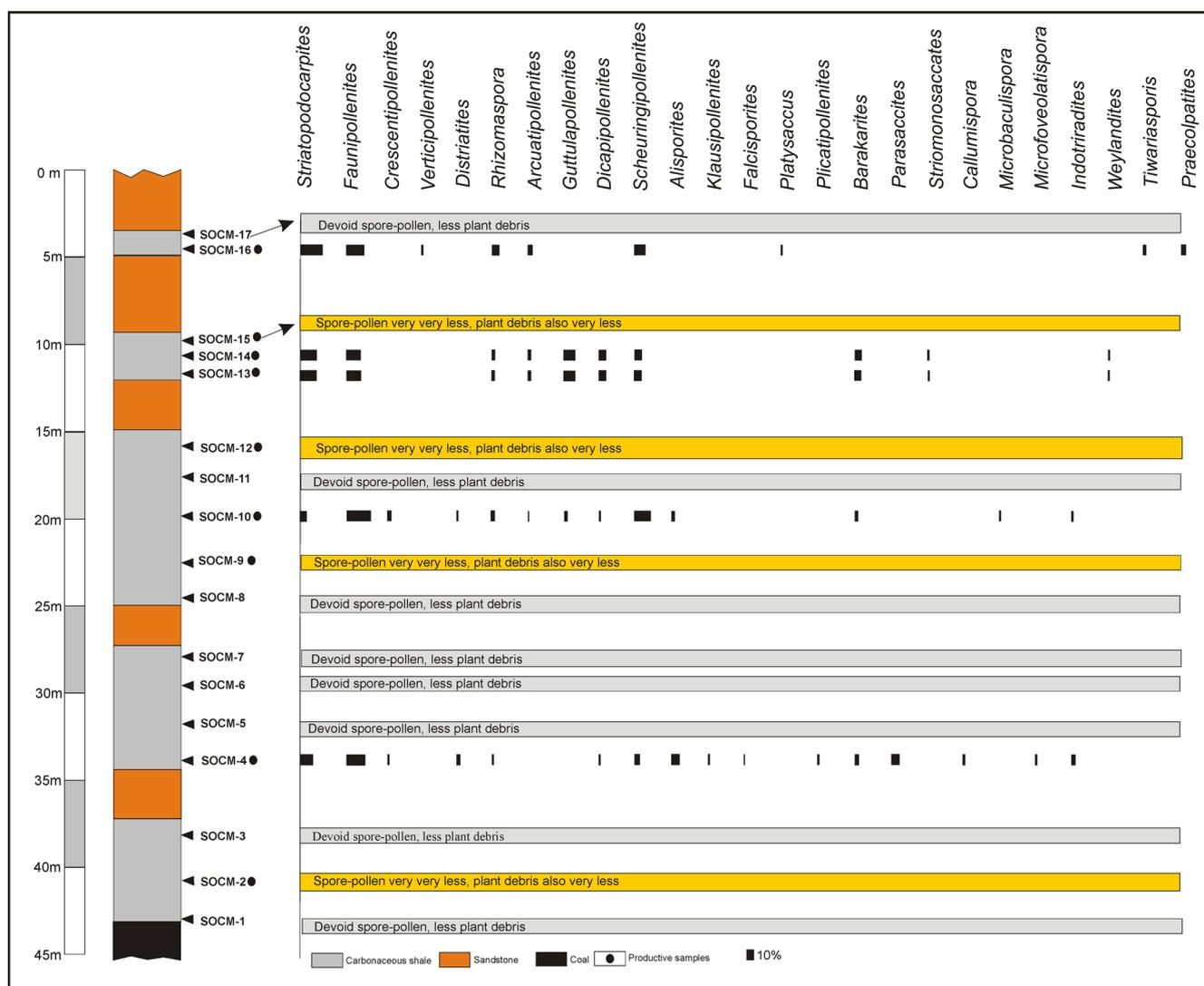


Fig. 2—A. Litholog and positions of samples. B. Distribution of palynotaxa in the Sawang OCM, East Bokaro Coalfield, Jharkhand; Dominant (>20%), subdominant (10–19%), common (5–9%), fair (2–4%) and poor (<2%).

sandstones and needle shale sediments. The major part of this basin to the east is occupied by Barakar Formation, which rests over the Talchir Formation and in other places the Barakar Formation rest directly on the Precambrian basement. This formation comprises coarse-grained arkosic sandstones, fine-grained laminated sandstones, grey shales, carbonaceous shales and coal seams. The crescentic outcrops of Barren Measures overlie the Barakar Formation and are exposed in the central, north-western and south-western regions, comprised of alternation of flaggy, cross-bedded, ripple-laminated ferruginous sandstones and shale beds. The successive overlying Raniganj Formation has a large spread along the base of the Lugu Hill and the formation is composed of medium to coarse-grained sandstones, shales and a few thin coal seams. The younger Panchet Formation occupies a vast area along the northern, eastern, and southern flanks of Lugu

Hill and is composed of fine-grained, greenish, micaceous sandstones and greenish shales, course-grained brownish-yellow, ferruginous sandstones and greenish chocolate sandy shales. The youngest strata in this coalfield, the Supra-Panchet Formation, consists of coarse clastics and rests over the Panchet Formation with an apparent angular unconformity, and this unconformable junction can be observed on the eastern face of the Lugu Hill. The composition of this formation is mostly coarse-grained ferruginous sandstones with lenses of pebbles. Interbedded within the sandstones are a few thin beds of red clays (Fig. 1).

## MATERIALS AND METHOD

The studied material comprises seventeen samples collected from the Sawang OCM section which is situated

Age	Formation	Lithology
Lower Cretaceous	Intrusive	Lamprophyre and dolerite dykes and sills
Upper Triassic	Supra-Panchet (Mahadeva ?)	Coarse-grained ferruginous sandstone, pebbly sandstone and red clay (600 m).
UNCONFORMITY		
Lower Triassic	Panchet	Greenish micaceous sandstone, buff fine-grained sandstone, red and green shale (500–600 m).
Upper Permian	Raniganj	Medium-to-coarse grained calcareous, sandstone, fine-grained greenish sandstone, grey shale, carbonaceous shale and thin coal seams (600 m).
Middle Permian	Barren Measures	Flaggy, fine-grained ferruginous sandstone micaceous sandy shale and black shale with siderite band (500 m).
Lower Permian	Barakar	Coarse-grained arkosic sandstone, fine-grained laminated sandstone, grey shale, carbonaceous shale and coal seams (900 m).
U–Carboniferous to L–Permian	Talchir	Tillite, greenish sandstones and needle shale
Pre–Cambrian	Pre–Cambrian	Granite gneisses, amphibolites and Mica schist

Table 1—Generalized stratigraphic succession of the East Bokaro Coalfield (after Raja Rao, 1987).

in the north-western part of the EBC. The succession is approximately of 45 m thick section and the lithofacies comprises mainly of carbonaceous shales, fine-grained sandstones and coal (Fig. 2A). 50 grams of each sample were taken and crushed (2–4 mm) and treated with 40% Hydrofluoric acid for 3–4 days to remove the silica content. Thereafter, the samples were washed thoroughly with distilled water to remove the acid content. The resultant residue was oxidized with concentrated Nitric acid and then treated with 10% Potassium Hydroxide solution. Five slides were prepared from each residue and the palynofossils were examined under standard light microscope (Olympus BX61 with DP–25 camera using Cell A software). Of the seventeen samples analysed, nine yielded pollen–spores which have been used for palynodating of the sediments (Fig. 2B).

### PALYNOLOGICAL ANALYSIS

Out of the nine yielding samples, five samples (i.e. SOCM 4, 10, 13, 14 and 16) were rich in palynofossils, while

four samples SOCM 2, 9, 12 and 15 were poor in palynofossils but rich in plant debris and amorphous organic matter. Besides, the remaining of samples (SOCM 1, 3, 5, 6, 7, 8, 11 and 17) were devoid of palynofossils and also poor in plant debris. The preservation of the palynofossils is variable within the samples, and recovery is low to moderate, light yellowish to dark brown, distorted, broken to fairly well-preserved (Pl. 1). The percentage frequency of the palynofossils is given in Figure 2B and indicated as dominant (more than 20%), subdominant (between 10–20%), common (between 5–9%), fair (between 2–4%) and poor (less than 2%) (Table 3).

The statistical analysis carried out in the studied section reveals that the pollen grains are more frequent compared to spores. Among the pollen, the representatives of striate bisaccate genera *Striatopodocarpites* (*S. labrus*, *S. ovalis*, *S. ovatus*, *S. subcircularis*) are the dominant taxa followed by *Faunipollenites varius*. The other associated common striate bisaccate taxa which characterise the assemblage include *Crescentipollenites fuscus*, *Verticypollenites gibbosus*, *Rhizomaspora* (*R. indica*, *R. triassica*), *Striatites varius* and

### PLATE 1



1. *Cyclogranisporites* sp.
2. *Callumispora* sp.
3. *Dictyotriletes invisus* Bharadwaj & Salujha (1964)
4. *Indotriradites korbaensis* Tiwari (1964)
5. *Indotriradites sparsus* Tiwari (1965)
6. *Parasaccites obscurus* Tiwari (1965)
7. *Scheuringipollenites maximus* (Hart) Tiwari (1973)
8. *Guttulapollenites hannonicus* Goubin (1965)
9. *Dicapipollenites* sp.
10. *Tiwariasporis novus* (Srivastava) Bharadwaj & Dwivedi (1981)
11. *Faunipollenites varius* Bharadwaj (1962)
12. *Striatopodocarpites labrus* Tiwari (1964)
13. *Crescentipollenites fuscus* (Bharadwaj) Bharadwaj *et al.*, (1974)
14. *Rhizomaspora triassica* Tiwari & Rana (1981)
15. *Striatites varius* Kar (1968)
16. *Platysaccus densicarpus* Anand–Prakash (1972)
17. *Falcisporites zapfei* Leschik emend. Klaus (1963)
18. *Alisporites indicus* Bharadwaj & Srivastava (1969)
19. *Klausipollenites schaubergii* Potonié & Klaus (1954)
20. *Arcuatipollenites pellucidus* (Goubin) Tiwari & Vijaya (1995)

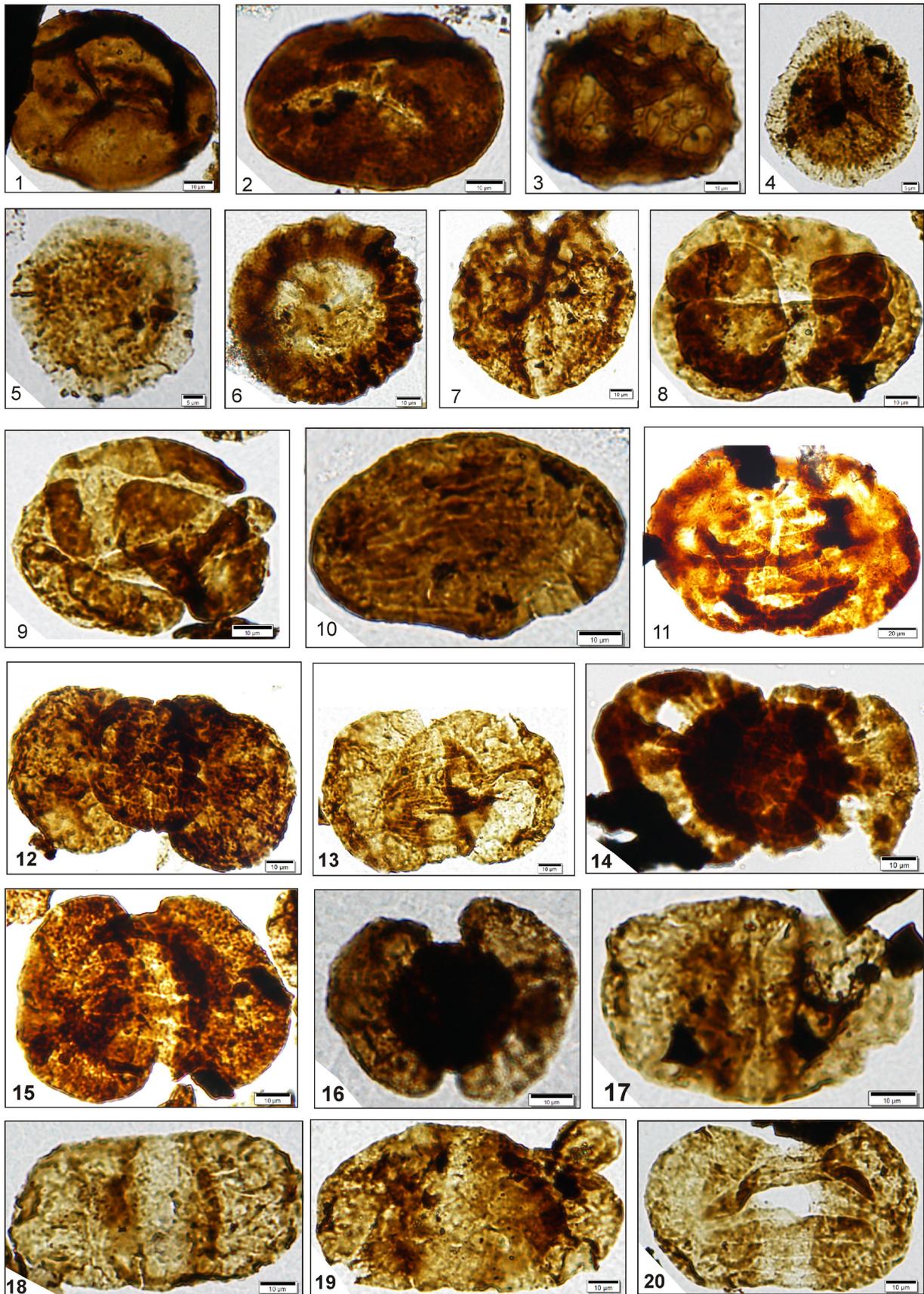


PLATE 1

*Distriatites* sp. Taeniate genera include *Arcuatipollenites* (*A. pellucidus* and *A. ovatus*), *Guttulapollenites hannonicus*, *Dicapipollenites* sp.; nonstriate bisaccate forms mainly include *Scheuringipollenites* (*S. barakarensis*, *S. tentula* and *S. maximus*) and others which are poor in frequencies, such as *Alisporites indicus*, *Falcisporites zapfei*, *Klausipollenites schaubegeri* and *Platysaccus densicarpus*. The monosaccates are poor in the assemblage and are represented by *Parasaccites obscurus*, *Plicatipollenites* sp., *Barakarites* sp. and *Striomonosaccites* sp. Inaperturate pollen are represented by *Weylandites* (*W. indicus*, *W. irregularis*), *Tiwariaspis novus* and *Praecolpatites* sp. The spores are meagre and represented by *Callumispora gretensis*, *Callumispora* sp., *Microfoveolatispora gondwanensis*, *Dictyotriletes invisus*, *Convertubisporites* and *Indotriradites* (*I. korbaensis*, *I. sparus*). The palynoassemblage mostly represent gymnosperm pollen in dominance, and pteridophytic spores are less in numbers.

The studied palynocomposition compares well with the *Striatopodocarpites–Faunipollenites* Palynozone–A (R–II B) in the Raniganj Formation of the Damodar Basin, which is dated as late Permian in age (Tiwari & Tripathi, 1992).

### CORRELATION

The Gondwana sequences of peninsular India exhibit different sedimentation patterns in each basin due to variable deposition in linear, fault bounded belts in which recurrent

uplift and subsidence at varying rates created different tectonic regimes (Jha *et al.*, 2014). Therefore, there are problems in inter-basinal correlations in the lithological context. However, they display broad similarities of palynoassemblages at the generic levels, thus favouring correlations within the Gondwana basins in peninsular India (Fig. 3, Table 2) and other Gondwana continents, viz. Australia, Antarctica, Africa and South America.

### Correlations with other Gondwana basins of peninsular India

The late Permian palynofossils of the present study are well correlated with other Gondwana basins of peninsular India, such as Damodar Basin (Bharadwaj & Tiwari, 1977; Rana & Tiwari, 1980; Singh & Tiwari, 1982; Tiwari & Singh, 1983; Tiwari & Tripathi, 1992; Srikanta Murthy, 2010; Srikanta Murthy *et al.*, 2010; Vijaya, 2011; Vijaya *et al.*, 2012; Srikanta Murthy *et al.*, 2014; 2016); Son–Mahanadi Basin (Srivastava *et al.*, 1977; Tiwari & Ram–Awatar, 1989; Tripathi & Bhattacharya, 2001; Srivastava & Kar, 2001; Kar, 2003; Kar & Srivastava, 2003; Ram–Awatar *et al.*, 2004); Satpura Basin (Bharadwaj *et al.*, 1978; Kumar, 1996; Srikanta Murthy *et al.*, 2013); Rajmahal Basin (Tripathi, 1986, 1989) and Wardha–Godavari basins (Srivastava & Jha, 1988, 1990, 1991, 1992, 1995; Jha & Srivastava, 1996; Jha *et al.*, 2007, 2014; Jha, 2008; Aggarwal, *et al.*, 2015; Jha & Aggarwal, 2015) in having similar

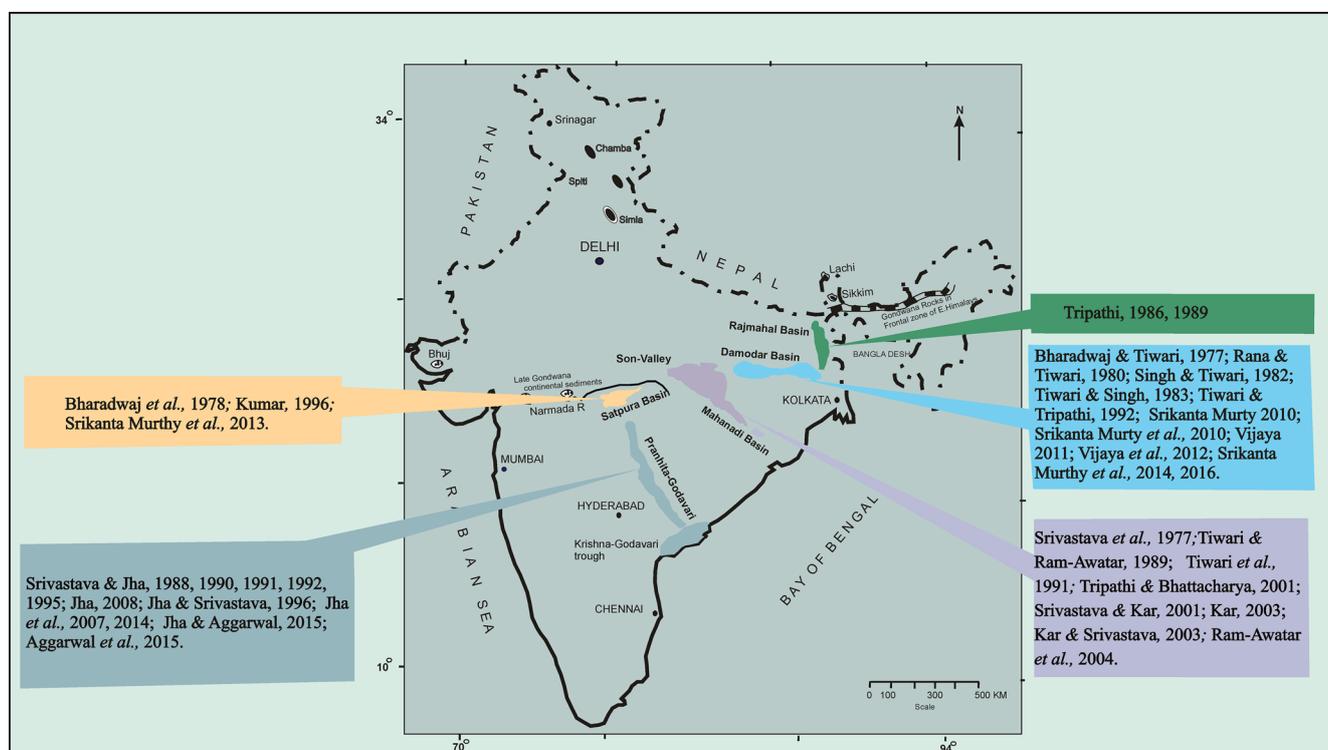


Fig. 3—Reports of Late Permian palynoassemblage from the Gondwana Basins of Peninsular India.

Gondwana Basins	Coalfields/Area	References
Damodar Basin	Karanpura Coalfield	Srikanta Murthy <i>et al.</i> , 2014.
	Raniganj Coalfield	Srikanta Murthy, 2010; Srikanta Murthy <i>et al.</i> , 2010; Vijaya, 2011; Bharadwaj & Tiwari, 1977; Tiwari & Singh, 1983; Singh & Tiwari, 1982; Rana & Tiwari, 1980; Tiwari <i>et al.</i> , 1992.
	East Bokaro Coalfield	Vijaya <i>et al.</i> , 2012; Srikanta Murthy <i>et al.</i> , 2016; Present study—Sawang OCM
Son–Mahanadi	Talchir Coalfield	Tripathi & Bhattacharya, 2001.
	Johilla Coalfield	Tiwari & Ram–Awatar, 1989.
	Sohagpur Coalfield	Ram–Awatar <i>et al.</i> , 2004;
	Tatapani–Ramkola	Kar, 2003; Kar & Srivastava, 2003; Srivastava & Kar, 2001; Srivastava <i>et al.</i> , 1977.
Satpura Basin	Shivapura Coal Mine	Srikanta Murthy <i>et al.</i> , 2013.
	Tamia Ghat Road	Kumar, 1996.
	Near Sukhtawa nala	Bharadwaj <i>et al.</i> , 1978.
Wardha–Godavari Basin	Kamptee Coalfield	Srivastava & Jha, 1988, Jha & Srivastava, 1996.
	Amavaram Area	Srivastava & Jha, 1991
	Mailaram Area	Srivastava & Jha, 1990,
	Budharam Area	Srivastava & Jha, 1995.
	Sattupalli Area	Srivastava & Jha, 1992.
	Chintalapudi	Jha, 2008.
	Gauridevpet Area	Jha <i>et al.</i> , 2014.
	Manuguru Area	Srivastava & Jha, 1992.
	Kachinapalli	Jha & Aggarwal, 2015.
	Mamakannu Area	Aggarwal <i>et al.</i> , 2015.
Rajmahal Basin	Birbhum Coalfield	Tripathi, 1986, 1989.

Table 2—Reports of Late Permian palynoassemblage from the Indian Peninsula.

palynoassemblages, such as dominance of *Striatopodocarpites* spp. and *Faunipollenites* spp. in association with other stratigraphically significant palynofossils, such as *Platysaccus*, *Crescentipollenites*, *Weylandites*, *Striomonosaccites*, *Barakarites*, *Alisporites*, *Falcisporites*, *Klausipollenites*, *Rhizomaspora*, *Arcuatipollenites*, *Microfoveolatispora*, *Dictyotriletes* and *Indotriradites*. Even though, these basins show a broad similarity, slight differences are also recorded from Sawang OCM in not having some palynofossils like *Corisaccites*, *Gondisporites*, *Lundbladispota*, *Distriatites*, *Brevitriletes*, *Horriditriletes*, *Densipollenites*, etc.

#### Correlation with the late Permian Gondwana counterparts

The late Permian assemblage of the present study is also correlatable with other Gondwana continents such as Antarctica (Balme & Playford, 1967; Kemp, 1973; Kyle & Schopf, 1982; Playford, 1990; Lindstrom, 1996); Australia

(Foster, 1979, 1982); Africa (Falcon, 1975; Anderson, 1977; Hankel, 1992; Modie & Le Herisse, 2009); South America (Marques–Toigo, 1991; Souza & Marques–Toigo, 2003, 2005; Souza, 2006; Gutierrez *et al.*, 2011) in having dominant striate bisaccate pollen grains mainly *Striatopodocarpites* and *Faunipollenites* (= *Protohaploxylinus*) and presence of some common taxa such as *Guttulapollenites*, *Lunatipollenites* (= *Arcuatipollenites*), *Scheuringipollenites*, *Alisporites* and *Klausipollenites*.

#### PALAEOENVIRONMENT

The palynological analysis from the Sawang OCM shows the dominance of striate bisaccates—*Striatopodocarpites*, *Faunipollenites* (= *Protohaploxylinus*), *Crescentipollenites*, *Verticypollenites*, *Rhizomaspora*, *Striatites*, and taeniate bisaccate *Arcuatipollenites* indicating the presence of conifers in the peat forming vegetation. Conifers are considered to be extra-basinal or hinterland elements, which typically

Sample No.	Palynocomposition
SOCM 4	<p>Dominance of striate bisaccate genera chiefly <i>Striatopodocarpites</i> (20–32%) followed by <i>Faunipollenites</i> (20–26%). Other stratigraphically significant taxa bisaccate, viz. <i>Crescentipollenites</i> (2–4%), <i>Verticypollenites</i> (2–3%), <i>Rhizomaspora</i> (2–10%), <i>Striatites</i> (2–5%) and <i>Distriatites</i> (0–2%). Taeniate genera include <i>Arcuatipollenites</i> (1–6%), <i>Guttulapollenites</i> (4–17%), <i>Dicapipollenites</i> (2–11%); nonstriate bisaccate forms mainly include <i>Scheuringipollenites</i> (7–20%) and others which are poor in frequencies, such as <i>Alisporites</i> (4–11%), <i>Falcisporites</i> (2–3%), <i>Klausipollenites</i> (0–1%) and <i>Platysaccus</i> (0–2%). The monosaccate are poor in the assemblage and are represented by <i>Parasaccites</i> (2–11%), <i>Barakarites</i> (4–7%), <i>Plicatipollenites</i> (0–2%) and <i>Striomonosaccites</i> (0–6%). <i>Inaperturopollenites</i> pollen are represented by <i>Weylandites</i> (2–5%), <i>Tiwariasporis</i> (0–4%) and <i>Precolpatites</i> (0–6%). The spores are meagre and represented by <i>Callumispora</i> (0–3%), <i>Microfoveolatispora</i> (0–2%), <i>Dictyotriletes</i> (0–4%), <i>Convertubisporites</i> (0–2%) and <i>Indotriradites</i> (0–2%).</p>
SOCM 10	
SOCM 13	
SOCM 14	
SOCM 16	

Table 3—Palynocomposition of Sawang Open Cast Mine.

show several adaptations for survival in drier habitats. Monosaccate forms are poor in percentage and low in diversity being represented by *Parasaccites*, *Plicatipollenites*, *Barakarites* and *Striomonosaccites*, reflecting the presence of Cordaites also in the peat forming flora. Cordaites pollen prefers mesophilous palaeoenvironment which is inhabited in well drained and low land substrates (Taylor & Taylor, 1993). Fragmentary presence of Cordaites suggests incursion of remnants of a parautochthonous seasonal dryland flora in the depositional environment (Jasper *et al.*, 2006). The nonstriate bisaccate pollen is represented by glossopterids such as *Scheuringipollenites*, *Alisporites*, *Falcisporites*, *Klausipollenites* and *Platysaccus*, indicating their prevalence in the peat forming vegetation. Glossopterids grew in mesophilous to xerophilous palaeoenvironment and flourished in lowland peats; while conifers survived in areas distant to the mires (Knoll & Nicklas, 1987). The trilete spores are relatively low in percentages and represented by Filicopsids (*Microbaculispora*, *Microfoveolatispora*, *Dictyotriletes*) and Lycopsids (*Indotriradites*) and are related to herbaceous and arborescent groups, which flourished in hygrophilous and mesophilous environments (Cazzulo–Klepzig *et al.*, 2005). The abundance conifer pollen in the present assemblage suggests the dominance of arborescent vegetation in the form of a forest swamp, probably in a small distant marginal part of the mire. The growth of herbaceous lycopsids and filicopsids probably favoured a flooding environment (Dimichele & Phillips, 1985). The overall palynological analysis suggests that the Sawang OCM palaeomire occupied inland areas of the basin and was deposited under fresh water environment.

### DISCUSSION

Three palynoassemblage zones have been identified by Tiwari and Tripathi, 1992 in the late Permian of the Damodar Basin namely, *Densipollenites indicus* Assemblage Zone or Zone VII, indicating Barren Measures Formation (Kulti Formation), *Gondisporites raniganjensis* Assemblage Zone or

Zone VIII representing lower part of Raniganj Formation and *Densipollenites magnicarpus* Assemblage Zone or Zone IX marking the upper part of Raniganj Formation. The sediments of the present study are palynologically dated to the lower part of Raniganj Formation, which is of late Permian age belongs to *Striatopodocarpites–Faunipollenites* zone–A (Tiwari & Tripathi, 1992). Other forms like *Crescentipollenites*, *Arcuatipollenites*, *Alisporites*, *Klausipollenites*, *Falcisporites* and *Guttulapollenites*, which are quantitatively less, also support this view. It can be concluded that the Sawang OCM section is dated as late Permian on the basis of above palynological evidences. Further, this palynoassemblage is compared well with previously known late Permian palynoassemblages from other Gondwana basins in peninsular India and also with gondwanan continents such as Australia, Africa, Antarctica and South America. The morphological characters, such as thin central body, diversity in striation and haploxylo-noid construction in the present palynocomposition indicate warm climate (Tiwari & Tripathi, 1987). On the basis of palynofossils composition, the present Sawang OCM succession represents a peat-forming community mainly composed of gymnosperms (conifers, glossopterids and cordaites) together with pteridophytes (lycopsids and filicopsids).

### CONCLUSIONS

- Late Permian age is proposed for the sediments from the Sawang open cast mine on the basis of palynological study.
- Palynological study also revealed that the peat forming vegetation mainly composed of gymnosperms represented by glossopterids, conifers, cordaites together with pteridophytes (lycopsids and filicopsids).
- The morphological studies in the present palynocomposition indicate warm climate during the deposition of sediments.

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